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Automated urban rainfall-runoff model generation with detailed land cover and flow routing

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19 Abstract

20 Constructing hydrological models for large urban areas is time consuming and laborious due 21 to the requirements for high-resolution data and fine model detail. An open-source algorithm 22 using adaptive subcatchments is proposed to automate Storm Water Management Model 23 (SWMM) construction. The algorithm merges areas with homogeneous land cover and 24 common outlet into larger subcatchments, while retaining small-scale details where land cover 25 or topography is more heterogeneous. The method was tested on an 85 ha urban catchment in 26 Helsinki, Finland. A model with adaptive subcatchments reproduced the observed discharge at 27 the catchment outlet with high model-performance indices emphasizing the strength of the 28 proposed method. Computation times of the adaptive model were substantially lower than those 29 of a corresponding model with uniformly sized high-resolution subcatchments. Given that 30 high-resolution land cover and topography data are available, the proposed algorithm provides 31 an advanced method for implementing SWMM models automatically even for large urban 32 catchments without substantial manual workload. Simultaneously, the high-resolution land cover details of the catchments can be maintained where they matter the most. 33

34 Keywords

35 SWMM, urban hydrology, stormwater, subcatchment delineation, automation, flow routing

37 **1. Introduction**

38 Urban areas are characterized by fragmented, mosaic land cover leading to altered hydrological 39 cycle when compared to natural areas. The changing landscape due to urbanization has impacts 40 on heat balance and evaporation (Whitford et al., 2001; Zhou et al., 2011), snow cover and 41 snowmelt (Bengtsson and Semádeni-Davies, 2011), infiltration and storm runoff generation 42 (Sillanpää and Koivusalo, 2015), and biodiversity (Pauleit et al., 2005) amongst other effects. 43 To understand the hydrology-related processes in urban areas, accurate description and 44 understanding of the land cover spatial configuration is crucial (Fletcher et al., 2013; Salvadore 45 et al., 2015).

46 The high-resolution description of catchment details is important in urban hydrological models 47 (Cantone and Schmidt, 2009). While for runoff volumes the impact of spatial resolution may 48 be modest or even negligible (Ghosh and Hellweger, 2012; Goldstein et al., 2016; Krebs et al., 49 2014; Park et al., 2008), the high detail in land cover description is particularly important for 50 accurate simulation of peak flow rates (Elliott et al., 2009; Ghosh and Hellweger, 2012; Krebs 51 et al., 2014). In addition to the increased accuracy of runoff simulations, describing 52 subcatchments in high-resolution as detailed units with homogeneous land cover simplifies the 53 model calibration procedure and narrows parameter ranges (Krebs et al., 2013; Sun et al., 54 2014).

In urban areas, impervious surfaces contribute the most to urban runoff and understanding their connection to surrounding areas and to the stormwater network is fundamental (Jacobson, 2011; Mejía and Moglen, 2010; Shuster et al., 2005). To accurately represent the flow routing between contributing surfaces in urban hydrological models, flow paths have to be described in detail requiring high-resolution data (Gironás et al., 2010; Rodriguez et al., 2013). The demand for high-resolution spatial descriptions of urban areas is also driven by the assessment of stormwater management systems, such as low-impact development and nature-based
solutions. These are often spatially distributed to individual surfaces or outlets of impervious
plots and need to be described in great spatial detail in models (Tuomela et al., 2019).

The US EPA Storm Water Management Model (SWMM) (Rossman, 2015) is a widely used open-source urban hydrological simulation model used for both event-based (e.g., Kong et al., 2017; Niemi et al., 2017) and long-term (e.g., Guan et al., 2015; Peleg et al., 2017; Taka et al., 2017) hydrological assessments in urban areas. Manual construction of high-resolution urban hydrological models, where each contributing surface is individually described (Krebs et al., 2014, 2013), is only feasible when the studied area is small. For larger areas, such as entire suburbs, automated methods are necessary to keep the task manageable.

71 Several tools have been proposed to facilitate the task of urban hydrological model 72 construction. Kertesz et al. (2007) developed a tool to compile and transfer subcatchment 73 information from ArcGIS Geographical Information System (GIS) to SWMM. Pina et al. 74 (2011) introduced an open-source tool inp.PINS for both creating SWMM input files directly 75 from GIS and for visualizing SWMM results in GIS, but the tool has since become deprecated. 76 Dongquan et al. (2009) presented a digital elevation model (DEM)-based automated batch 77 process for subcatchment discretization in ArcGIS without accounting for different land covers 78 within subcatchments and tested the results with SWMM. In addition, several commercial 79 modelling packages built around the EPA SWMM computational engine exist (e.g., 80 InfoSWMM and XPSWMM by Innovyze, PCSWMM by Computational Hydraulics 81 International) to aid the modeler e.g., by incorporating superior GIS capabilities over the 82 standard EPA SWMM user interface or by allowing integrated 1D-2D modelling using 2D surface flow descriptions. Nevertheless, even with commercial packages, subcatchment 83 84 delineation and routing of water between subcatchments and into the stormwater network are still largely manual tasks. Clearly, there is still room for improvement as none of the tools
facilitate automatic model building while retaining the detailed land cover characteristics of
the urban environment.

Easily available, remotely sensed high-resolution data on topography and land cover are abundant (Bates, 2004; French, 2003; Tarolli et al., 2013). This shifts the focus in hydrological modelling from describing the environment in as much detail as possible into such models where simulation times remain feasible (Bates, 2012; Sampson et al., 2012). High-resolution description of land cover and flow paths becomes important in urban environments and requires a modeler to find a balance between necessary level of detail and acceptable computational burden.

95 Describing a sizeable urban area in high-resolution often results in unfeasibly long simulation times necessitating for a method to aggregate adjacent surfaces into larger computational units. 96 97 Warsta et al. (2017) proposed an automatic method for building SWMM models in a manner 98 where each DEM/land cover raster cell corresponded to one subcatchment. The method also 99 allowed combining individual grid cells into larger rectangular subcatchments in a rudimentary 100 manner. While this decreased computation times and facilitated application to large urban 101 catchments (Rautiainen, 2016), averaging catchment properties (e.g., elevations) while 102 combining grid cells led to problems with surface runoff routing and to a loss of fine-scale 103 detail in describing land cover and topography.

To tackle the challenge of automatically constructing SWMM models in high-resolution while minimizing the computational burden, the main objective of this paper was to propose a new open-source algorithm to automate SWMM model construction. Following the requirements for accurate flow path description and homogenous subcatchment land cover, the proposed algorithm automatically discretizes the studied area in an adaptive manner based on land cover

and flow routing using high-resolution land cover and DEM data. The result is a SWMM model
with a minimum number of subcatchments where each subcatchment is covered by a single
land cover type.

The performance of the new discretization method was demonstrated by comparison against a uniform discretization scheme where each raster cell corresponds to one subcatchment. Simulation results were also compared against field measurements to validate model performance.

116 2. Materials

The studied Länsi-Pakila catchment (Fig. 1) is an 85 ha urban area in Helsinki, Finland. Helsinki has a boreal climate with a mean annual air temperature of 5.9°C and mean annual precipitation of 655 mm, with most of the rainfall falling in late summer and early autumn (Pirinen et al., 2012).

121

[FIGURE 1]

The Länsi-Pakila catchment is a medium-density residential area characterized by detached houses. The area is relatively green, with vegetation covering 53.5%, asphalt 27.5%, and roofs 13.5% of the area (Table 1), resulting in a total imperviousness of 43%. The area is prone to stormwater flooding (Raukola, 2012). Länsi-Pakila is subject to urban development and faces a risk of more severe urban flooding in the future unless due attention is paid to stormwater management.

128

[TABLE 1]

129 An openly available 1×1 m² DEM from the City of Helsinki was used for catchment 130 delineation. The catchment land cover description was based on the openly available land cover

classification data from the Helsinki Region Environmental Services Authority HSY. As is
often the case with urban runoff studies, several site visits were required to complement the
scattered stormwater network information available from the network map.

Rainfall was measured at the Länsi-Pakila catchment during summer 2017 using three colocated fully automatic tipping-bucket rain gauges (ECRN-100 High Resolution Rain Gauge) with 0.2 mm tip size and 1 min temporal resolution. The gauges were located on top of a lowrise nursing home building to keep them safe from vandalism and to minimize obstruction from the urban surroundings (Fig. 1). Daily air temperature and wind speed data were available from the Finnish Meteorological Institute's weather station in Kumpula, Helsinki, approximately 5 km south-east from the catchment.

141 Catchment discharge information was obtained by measuring water level and flow velocity at 142 the catchment outfall (Fig. 1) in an 800 mm concrete pipe using a Starflow Ultrasonic Doppler 143 Instrument Model 6526. The time resolution of the discharge measurements was 1 min. The 144 instantaneous discharge measurements caused velocity fluctuations, which were smoothed 145 using a 5 min central moving average in further data preparation.

146 **3. Methods**

147 **3.1. Adaptive subcatchment discretization**

The proposed subcatchment discretization algorithm extends the automatic SWMM model construction tool introduced by Warsta et al. (2017). They divided the investigated area into subcatchments using a uniform computation grid with a desired spatial resolution. The grid cells were then connected to each other and into the stormwater network. Following Krebs et al. (2014), the generated subcatchments were small enough to be hypothesized to consist of a single homogenous land cover type, e.g., a green area, a rooftop, or a paved road. This

154	simplifies the calibration of the resulting SWMM model by allowing parameters to be linked
155	to distinct surface types. In the proposed algorithm, the subcatchments are still assumed to be
156	of a homogeneous land cover type, but their sizes depend on the underlying land cover and
157	flow routing. This reduces the number of subcatchments greatly in areas where land cover and
158	topography are uniform.

- 159 The proposed algorithm proceeds as follows:
- A model with a uniform computation grid where each grid cell corresponds to one
 subcatchment is created for the studied area.
- 162 2. Cells with open storm sewer nodes are initially saved as a set of one-cell-sized
 163 subcatchments. SWMM parameters for each subcatchment are adopted from the node
 164 cell.
- 165 3. Subcatchments are processed one by one.
- 166 4. All adjacent upstream cells routed into the currently processed subcatchment are listed:
- a. If an upstream cell has the same land cover as the currently processed
 subcatchment, the cell is merged to the subcatchment. Subcatchment parameters
 (area, elevation, slope) are updated.
- b. If an upstream cell has different land cover than the currently processed
 subcatchment, a new subcatchment is created. Subcatchment properties are
 copied from the underlying cell for the newly created subcatchment.
 Downstream subcatchment is set as the outlet of the new subcatchment.
- 5. Subcatchments and upstream cells are traversed until all cells contributing to any of theopen storm sewer nodes are processed.
- 176 6. Neighbouring roof cells sharing their outlet are merged together to form a new
 177 subcatchment. Subcatchment parameters (area, elevation, slope) are computed for each

merged roof subcatchment. Depending on the given land cover class for the roof, theroof subcatchments are routed either

a. to the nearest adjacent (non-roof) subcatchment (disconnected roof in Fig. 1) orb. directly to the nearest storm sewer node (connected roof).

182 Subcatchment cells with an open storm sewer node are connected into the stormwater network. 183 Furthermore, a 3×3 cell area surrounding each open storm sewer node is used as a collecting 184 area for the node to account for errors in flow routing resulting from inaccuracies in the DEM. Cells not contributing to any downstream subcatchment or storm sewer node, e.g., at the 185 186 borders of the study area, are disregarded from further analysis. Subcatchment area is the 187 combined area of the contributing cells, whereas subcatchment elevation and slope are assigned 188 the average of the contributing cells. The flow width (FW) parameter for adaptive 189 subcatchments is approximated after Krebs et al. (2014) as

 $FW = k\sqrt{A} \tag{1}$

191 where *A* is the subcatchment area and k = 0.7.

192 The proposed algorithm requires three raster files with equal dimensions and resolution as 193 inputs: a land cover raster where different land cover classes are identified with integers, a DEM raster depicting the topography of the studied area, and a flow direction raster with 194 195 integers from 1 (north-east) to 8 (north) indicating the direction of flow from each raster cell. It is assumed that all cells in the flow direction raster are routed, i.e., there are no pits. Other 196 input files to the tool are identical to those reported by Warsta et al. (2017), consisting of 197 198 geometry files for the stormwater network, a file relating the land cover raster indices to 199 SWMM subcatchment parameters, and various settings files. Given the input files, the tool produces a SWMM input file (.inp) ready to be used in simulations and a set of GIS-compatible
files that facilitate visualization of model set-up in GIS software.

202 **3.2. Model implementations**

Two models for the Länsi-Pakila catchment were created; a model where all subcatchments are rectangular and have dimensions of 1×1 m² (referred to as *1x1*) and a model with adaptive subcatchments (*adap*) according to the proposed algorithm. The *1x1* model acts as the reference model for assessing the subcatchment discretization impact on simulation performance.

207 Six largest rainfall-runoff events from the Länsi-Pakila catchment were selected for analysis 208 (Table 2). Three of the events were used for model calibration and the remaining three were 209 validation events. Earlier, Sillanpää and Koivusalo (2014, 2015) showed a difference in urban 210 runoff response between minor and major storms due to the runoff-contributing area expanding 211 from impervious to pervious areas during major storms. Because the storm size may affect 212 model parameterization, and because the main interest was in potential urban flood-producing 213 events, the selected events were all major storms (rainfall accumulation >17 mm). The 214 threshold defining a major storm corresponds with the rainfall threshold set by Sillanpää and 215 Koivusalo (2014) and Guan et al. (2016) for similar climate and catchment conditions.

216

[TABLE 2]

Land cover in both 1x1 and adap was represented with 6 classes. All model parameter values except for infiltration parameters corresponded to those used by Warsta et al. (2017) and Krebs et al. (2014) for similar urban catchments in Finland (Appendix A). Initial tests showed SWMM to be sensitive to the Green-Ampt infiltration model parameters, i.e., the suction head (ψ_s) , the saturated hydraulic conductivity (K_s) , and the maximum soil moisture deficit (θ_{dmax}) , and therefore these parameters were selected for the model calibration. Infiltration parameters of the underlying soil were uniform for all land cover types and PEST (v.13) software (Doherty, 2016) with Tikhonov regularization was used to calibrate the *adap* model. The parameters for loamy sand from Rawls et al. (1992) were used as initial values. The calibration was conducted by minimizing the sum of squared errors of simulated flow against observed flow for time steps when observed flow exceeded a threshold of 7 - 15 l/s depending on the event. The threshold was selected due to the focus on potential flood-producing events and the desire to match peak flows rather than base flow. The same calibrated parameters were then used for the *lx1* model.

The flow direction raster was created from the DEM. As a pre-processing step, the stormwater network information was integrated into the DEM using "stream burning" (e.g., Saunders, 1999) to ensure maximum collecting area for the catchment. Subsequently, the original DEM without network burning from the corresponding area was used to produce the flow direction raster using r.watershed tool from GRASS GIS (GRASS Development Team, 2017). Cells with land cover classified as buildings were set to block the overland flow and cells with open storm sewer nodes were set to collect the water.

The model performance was evaluated against observations using the Nash-Sutcliffe efficiency *NSE* (-) (Nash and Sutcliffe, 1970), volume error (relative bias) *VE* (%), and peak flow error *PFE* (%).

240
$$NSE = 1 - \frac{\sum_{t} (Q_{o,t} - Q_{s,t})^2}{\sum_{t} (Q_{o,t} - \overline{Q}_{o})^2}$$
(2)

241
$$VE = 100 \frac{V_s - V_o}{V_o}$$
 (3)

242
$$PFE = 100 \frac{Q_{s,max} - Q_{o,max}}{Q_{o,max}}$$
(4)

where $Q_{o,t}$ and $Q_{s,t}$ are the observed and simulated discharge (l/s), respectively, at time t, $\overline{Q_o}$ is the average observed discharge (l/s) during an event, V_o and V_s are the observed and simulated flow volumes (m³), respectively, during an event, and $Q_{o,max}$ and $Q_{s,max}$ are the observed and simulated maximum discharges (l/s), respectively. The performance of *adap* was evaluated against lxl using the Pearson correlation coefficient r (-):

248
$$r = \frac{\sum_{t} (Q_{1,t} - \overline{Q_1})(Q_{a,t} - \overline{Q_a})}{\sqrt{\sum_{t} (Q_{1,t} - \overline{Q_1})^2} \sqrt{\sum_{t} (Q_{a,t} - \overline{Q_a})^2}}$$
(5)

where $Q_{1,t}$ and $Q_{a,t}$ are the simulated discharges (1/s) at time *t* from the *1x1* and the *adap* models, respectively, and $\overline{Q_1}$ and $\overline{Q_a}$ are the average simulated discharges (1/s) during an event using the *1x1* and the *adap* models, respectively. In addition, the performance was evaluated using volume difference *VD* (%) and the peak flow difference *PFD* (%) by substituting *1x1* and *adap* for observed and simulated in Eqs. (3) and (4), respectively.

254 **4. Results**

255 4.1. Adaptive subcatchment discretization

256 Table 3 presents the subcatchment statistics for *adap* and *1x1* subcatchment discretizations. The *adap* model resulted in only 10% (82 554) of the number of uniform $1 \times 1 \text{ m}^2$ 257 subcatchments in 1x1 (848 258). The subcatchment sizes for adap ranged up to 9 322 m² with 258 259 a mean size of 10.3 m². As both *adap* and 1x1 share the same input DEM data, the mean 260 subcatchment elevation and slope are equal. Differences in the range of subcatchment 261 elevations and slopes are due to some individual raster cell subcatchments of 1x1 being merged in adap. The maximum subcatchment slope of over 400% is explained by local errors in the 262 DEM. For 1x1, with all the subcatchments having constant dimensions of 1×1 m², the 263 264 subcatchment flow width is always either 1 or 0.7 m depending on whether the flow is perpendicular or diagonal from the cell. For *adap* subcatchments, the flow width was computed using Eq. (1) resulting in a larger range of flow widths for individual subcatchments. The average subcatchment flow width was however similar for both models; 0.9 and 1.4 m for 1x1and *adap*, respectively.

269

[TABLE 3]

270 Creating the *adap* models using the proposed algorithm consumed 30% more time (29.3 min) 271 than creating the 1x1 model (22.6 min). The reduction in the number of subcatchments led to 272 a corresponding reduction in SWMM computation time; the average computation time of an 273 *adap* SWMM model was 10% of the corresponding 1x1 model computation time for the 274 calibration and validation events (Fig. 2).

275

[FIGURE 2]

Fig. 3 shows the reduction in the number of subcatchments and the effect of combining 276 277 individual cells into larger subcatchments on flow routes when moving from 1x1 to adap. The 278 routing from one land cover type to another follows the same paths in *adap* and in 1x1. 279 However, in *adap* the number of subcatchments is substantially lower as cells with the same 280 land cover type and sharing a common flow path have been merged. Note that the flow routing 281 in Fig. 3 is displayed between subcatchment mass centers, creating an illusion that not all *adap* subcatchments are routed when in fact some subcatchment mass centers are situated outside 282 283 the subcatchment and/or the figure.

284

[FIGURE 3]

285 **4.2. Model calibration and validation**

286 Fig. 4 illustrates the *adap* and 1x1 simulation results for the calibration and validation events, 287 while Table 4 presents the corresponding performance statistics (Eqs. 2-5). The *adap* model 288 performed well for both calibration and validation events. For all events except v1, NSE coefficients exceeded 0.90 indicating "very good" adap model performance, and NSE of 0.8 289 290 for event v1 still indicated "good" model performance according to the recommended model-291 performance classes by Ritter and Muñoz-Carpena (2013). Regardless of the event, simulations 292 from *adap* tended to slightly underestimate flow volumes with the underestimate somewhat 293 larger for the validation events (average VE - 13.1%) than for the calibration events (-7.1%). 294 This was partly explained by a more rapid return to pre-storm flow levels in the simulations 295 than in the observed data, although some of the high flows were also underestimated. Peak 296 levels were mainly well captured, with *PFE* less than 10% except for events c3 (*PFE* -11.3%) and v3 (19.3%). Only for event v3 did *adap* overestimate the peak flow. 297

298

[FIGURE 4]

299

[TABLE 4]

Performance of the *1x1* model in terms of *NSE* varied from "acceptable" for events v1 and c1 to "very good" for c2 and c3 (Table 4). Unlike in *adap* simulations where volume error was negative for all events, in *1x1* the flow volume was overestimated for all events with *VE* ranging from 5.1% for v1 to 22.9% for c1. For events c1 and v3, *1x1* overestimated the peak flow by roughly 20%. However, for the other events the peak flow was accurately simulated.

The statistics between *adap* and IxI (Table 4) highlight the similar reaction of both models to rainfall events, as demonstrated by the correlation coefficients between *adap* and IxIapproaching unity. Still, *adap* constantly produced 15 – 25% lower flow volumes than IxI.

Although in absolute terms *VE* of both models was similar, *adap* on average underestimated observed flow volumes by 10.1% while *1x1* overestimated by 11.6%. The *adap* model generally predicted roughly 5% lower peak flows than *1x1* with the maximum *PFD* being -17.4% for event c1. For this event, the difference was almost entirely due to discharge overestimation of *1x1*.

The differences in peak flows and flow volumes are consistent with catchment mass balance differences between *adap* and 1x1 (Table 5). In each event, share of surface runoff was less for *adap* than for 1x1, whereas infiltration was greater for *adap* than for 1x1. In addition, *adap* produced slightly less evaporation than 1x1 while final stored water volume was slightly larger in *adap* than in 1x1.

318

[TABLE 5]

319 **5. Discussion**

320 Automated DEM-based methods for SWMM subcatchment generation have been proposed before, but they differ from the current algorithm. Dongquan et al. (2009) aimed for a low 321 322 number of computational units by using a high-resolution 2×2 m² DEM but combining all 323 cells belonging to the same drainage basin into subcatchments regardless of land cover or flow 324 routing details. Their approach resulted in 113 subcatchments for a 13.65 ha study area, vielding a hundred times coarser average subcatchment size of 1 200 m² than in the *adap* model 325 here (10.3 m²). Warsta et al. (2017) described the catchments in fine detail but because each 326 327 cell was considered an individual subcatchment, computation times were long with a large 328 number of redundant cells in areas with a homogenous land cover type. This is analogous to 329 the 1x1 simulations here. In the proposed adaptive algorithm, more subcatchments are 330 generated in areas where either land cover or flow routes are heterogeneous whereas in more

331 homogenous areas the subcatchments are allowed to have a larger size. The rudimentary grid 332 cell aggregation procedure of Warsta et al. (2017) yielded shorter computation times, but 333 changed the land cover and flow routing patterns in the catchment. This resulted in a moderate 334 reduction of simulated peak flows and flow volumes. Here, the computation time was greatly reduced as *adap* simulations took on average only 10% of the corresponding 1x1 simulation 335 336 time, while both *adap* and *1x1* produced good simulation results. The smaller computational 337 burden associated with the proposed algorithm allows model construction for large urban 338 catchments.

339 Because no manually constructed SWMM models exist for the studied catchment, direct 340 comparison of computation times to a corresponding manual model was not possible. However, 341 a crude estimate of a roughly five-fold increase in computation time between a manual model 342 and a corresponding *adap* model was approximated by comparing models from earlier studies. 343 In the work of Niemi et al. (2019), the proposed algorithm was used to create SWMM models 344 with adaptive subcatchments for three small urban catchments (5.87, 6.63, and 12.59 ha 345 catchment areas) in Lahti, Finland. Earlier, Krebs et al. (2014) manually constructed high-346 resolution SWMM models for the same catchments. The adap models in Lahti had, on average, 347 14.7 times the number of subcatchments when compared to the corresponding manual models, 348 and required, on average, 4.8 times as long to compute.

The adaptive subcatchment discretization algorithm retains the high spatial resolution of the input DEM and land cover data where necessary, but creates larger subcatchments where such spatial detail is not crucial. This allows for an accurate spatial representation of land cover, deemed important by Cantone and Schmidt (2009) and Petrucci and Bonhomme (2014). However, it also relieves the computational burden that can become excessive with a uniformly high spatial resolution model. Given that input land cover data are in raster format, the developed algorithm retains the land cover description from the input data. Otherwise, the
accuracy of the land cover description depends on the rasterization of non-raster-format input
data.

358 When building stormwater models manually, surfaces are usually assumed to drain entirely 359 into a single inlet node unless there is a compelling reason to resolve the routing in other way. 360 Therefore, an entire impervious surface, such as a parking lot, may be routed into a single inlet 361 although actual topography-driven flow paths would drain a part of the area to adjacent yards. 362 In the method of Warsta et al. (2017), routing of pit cells depended on their location in either 363 pervious or impervious areas. All water routed into pits residing in pervious areas was 364 infiltrated whereas water routed into pits in impervious areas was routed directly to the nearest 365 storm sewer node. As a result, some areas did not contribute to the catchment runoff as the 366 water had been infiltrated into a pit along its flow path. On the other hand, contribution of other 367 areas was unduly exaggerated as flow from them was routed directly to the stormwater 368 network. The proposed new algorithm allows the water to follow topography-driven flow 369 paths, and the use of a depressionless DEM ensures that water is routed through local pits. This 370 refined routing, compared to Warsta et al. (2017), should offer better runoff predictions during 371 major storms when pervious surfaces get saturated and start to convey runoff (Sillanpää and 372 Koivusalo, 2014; Yao et al., 2016).

Both *adap* and *1x1* models appropriately reproduced the observed runoff at the studied catchment. The slight underestimation of flow volumes by *adap* was expected, as the underestimation by SWMM in simulating hydrograph tails and low flows is commonly encountered (e.g., Guan et al., 2015, 2016). This behaviour was accentuated by calibration of *adap* focusing on high flows in lieu of low flows to more accurately simulate potential urban flood-producing events. 379 Because the model parameters in *adap* and 1x1 were identical, differences in simulated flow 380 volumes and peak flows between the implementations are explained by those model 381 characteristics that were different: flow width, subcatchment slope, and subcatchment area. As 382 these variables appear in the Manning equation SWMM uses to express the surface runoff for 383 each computation time step, the dynamics of runoff production are altered when the parameters 384 change. More importantly, the volume of infiltrated water within a subcatchment depends on 385 its size. As the volume of the infiltrated water is the product of the area and infiltration depth, 386 a larger subcatchment can infiltrate more water than an equally parameterized but a smaller 387 subcatchment. In *adap*, the average subcatchment size was larger than in 1x1 and the likelihood 388 of runon being completely or mostly infiltrated was larger. It is noteworthy, that the 389 dependence of infiltration volume on subcatchment size involves all SWMM models, 390 regardless of their construction procedure. The matter concerns especially models that treat 391 infiltration as a loss from the system without consideration of the storage capacity of the 392 underlying ground.

393 Adjusting only the infiltration parameters while taking other input parameters from Warsta et 394 al. (2017) and Krebs et al. (2014) was sufficient to yield a well-performing model, with relative uncertainty variance reductions of 0.61, 0.99, and 0.96 for ψ_s , K_s , and θ_{dmax} respectively. 395 396 These results are in line with earlier research suggesting that extensive calibration of a 397 hydrological model may be unnecessary if representative parameter sets are available from similar catchments (Bárdossy, 2007; Gao et al., 2015; Kokkonen et al., 2003; Krebs et al., 398 399 2016). The results also support the findings of Petrucci and Bonhomme (2014) stating that an 400 uncalibrated SWMM model may perform comparably to a calibrated model as long as land 401 cover is described accurately. The slightly less accurate discharge simulations from 1x1 than 402 adap were because infiltration parameters were calibrated using adap and applied to 1x1

403 without further calibration. However, had IxI also been calibrated the differences between the 404 models would likely be smaller.

405 SWMM performance is often found to be sensitive to the subcatchment flow width parameter 406 (Niazi et al., 2017). Here, the sensitivity to FW was assessed by evaluating the performance of 407 adap in event c1 with calibrated infiltration parameter values while allowing coefficient k in 408 FW (Eq. 1) to vary from 0.3 to 1.1 in steps of 0.2. The results showed *adap* to be rather 409 insensitive to FW (NSE variation between 0.91 - 0.94, PFE -1.36% - 1.43%, and VE -3.99%410 -4.44%), justifying the decision to use Eq. (1) with k = 0.7 to describe the flow width in this 411 study. However, due to the often encountered importance of proper flow width 412 parameterization in SWMM modelling, and the possibility to calculate it explicitly in the 413 proposed algorithm that traverses through raster cells, this should be considered as one of the 414 first improvements to the presented algorithm.

415 **6.** Conclusions

416 This study presented a new algorithm for automating SWMM model construction with a novel 417 solution to delineate subcatchments based on shared land cover and outlet. The algorithm 418 creates subcatchments adaptively by merging small subcatchments having homogeneous land 419 cover and common outlet into larger areas while retaining small-scale details where land cover 420 is heterogeneous. While pre-processing the input files for the proposed tool is convenient to 421 perform in a GIS software, the proposed tool itself is platform-independent, open-source, and 422 not tied to any specific GIS software. The tool facilitates urban hydrological assessments by 423 substantially reducing the required manual workload.

424 Based on the results obtained in this study, the following conclusions were drawn:

The proposed algorithm facilitates rapid model construction even for large urban areas 425 426 while retaining the high-resolution details where necessary. SWMM simulation results obtained using the proposed algorithm matched well with 427 428 catchment discharge observations. 429 Use of adaptive subcatchments resulted in a substantial reduction in the computational 430 burden while yielding similar simulation results to a model having a uniformly high-431 resolution subcatchment delineation. 432 • Good model performance obtained with an existing parameter set from similar 433 catchments conditions, adjusting only infiltration parameters, is encouraging regarding 434 stormwater predictions in ungauged urban areas. 435 The main limitation of the proposed tool is the requirement for high-resolution and high-quality land cover and DEM data. 436

437 **7. Appendix A**

Table A1 presents the land cover parameter values used in *adap* and *1x1* adopted from Warsta
et al. (2017) and Krebs et al. (2014) for similar urban catchments in Finland. The Green-Ampt
infiltration parameters are based on model calibration.

441

[TABLE A1]

442 8. Acknowledgements

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603 Tables

Land cover	Fraction (%)
Vegetation	53.50
Asphalt	27.51
Connected roofs	7.87
Disconnected roofs	5.66
Sand and gravel	5.17
Water	0.23
Rock outcrops	0.06

Table 2. Summary statistics of the studied rainfall-runoff events. Events c1-c3 were calibration

607 and v1-v3 va	alidation events.
------------------	-------------------

Event code	Date	Event duration	Rainfall depth	Peak rain intensity	Flow volume	Peak flow
		(h)	(mm)	(mm/min)	(m ³)	(l/s)
c1	6 Jun 2017	20	30.0	0.6	4 321	540
c2	2 Aug 2017	9	17.6	0.4	2 567	368
c3	9 Sep 2017	13	19.8	0.4	2 882	364
v1	12 Jun 2017	19	23.2	0.2	3 263	309
v2	4 Aug 2017	13	31.4	1.0	5 368	509
v3	12 Sep 2017	7	23.6	1.0	4 035	615

609 Table 3. Subcatchment statistics in *adap* (82 554 subcatchments) and *1x1* (848 258

610 subcatchments) SWMM models.

Statistic		adap			1x1		
	min	mean	max	min	mean	max	
Area (m ²)	1.0	10.3	9 322.0	1.0	1.0	1.0	
Elevation (m.a.s.l.)	19.1	27.5	45.8	13.4	27.7	45.8	
Flow width (m)	0.7	1.4	67.6	0.7	0.9	1.0	
Slope (%)	0.2	5.5	417.7	0.1	5.5	464.0	

- **Table 4.** Performance statistics of the *adap* and the *1x1* model simulation results against
- 613 observations (*obs*) and of the *adap* against the 1x1 model simulation results for the calibration

Event	adap vs. obs			<i>1x1</i> vs. <i>obs</i>			<i>adap</i> vs. 1x1		
Event	NSE(-)	VE (%)	PFE (%)	NSE (-)	VE (%)	PFE (%)	r (-)	VD (%)	PFD (%)
c1	0.92	-4.1	0.7	0.74	22.9	21.9	0.97	-22.0	-17.4
c2	0.97	-9.4	-2.2	0.94	9.8	7.8	0.99	-17.5	-9.3
c3	0.96	-7.7	-11.3	0.97	8.8	-3.4	0.99	-15.2	-8.1
v1	0.80	-9.6	-5.2	0.70	5.1	-0.1	0.99	-14.0	-5.1
v2	0.92	-16.3	-3.5	0.89	6.2	0.9	0.97	-21.2	-4.3
v3	0.91	-13.3	19.3	0.84	16.4	21.3	0.95	-25.5	-1.6

614 (c1-c3) and the validation (v1-v3) events.

c^2 17.600 c^3 19.800
0.247 16.174 3.173
16.174 3.173
0 107
110
72 700
0016
0.230 13.713
1000

617 **Table 5.** Mass balance statistics of the *adap* and the 1x1 model simulation results for the 618 calibration (c1-c3) and the validation (v1-v3) events.

Note: P = precipitation; E = evaporation; I = infiltration; R = surface runoff; S = final storage; CE = continuity error.

620 **Table A1.** SWMM parameter values for six surface classes and three stormwater network

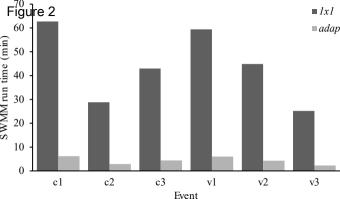
Surface type	I (%)	D (mm)	n (-) k	K _s (mm/h) ^a	$\psi_s~(\mathrm{mm})^{\mathrm{a}}$	θ_{dmax} (-) ^a
Asphalt	100	0.42	0.011	24.965	55.832	0.350
Rock outcrop	100	2.49	0.030	24.965	55.832	0.350
Roof	100	0.87	0.012	24.965	55.832	0.350
Sand, gravel	33	2.49	0.030	24.965	55.832	0.350
Vegetation	0	4.22	0.238	24.965	55.832	0.350
Water	100	0.10	0.011	24.965	55.832	0.350
Concrete pipe	-	-	0.015	-	-	-
PVC pipe	-	-	0.011	-	-	-
Open channel	-	-	0.049	-	-	-

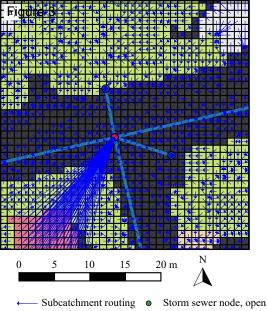
621 classes in *adap* and *1x1* models. Adopted from Warsta et al. (2017) and Krebs et al. (2014).

622 Note: I = imperviousness; D = depression storage; n = Manning's roughness; $K_s =$ saturated hydraulic conductivity; $\psi_s =$

623 suction head; θ_{dmax} = maximum moisture deficit; ^a calibrated parameter.

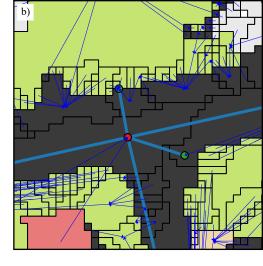






•

Stormwater network



Storm sewer node, closed

Rooftop, connected Rooftop, disconnected

Sand and gravel Asphalt Vegetation

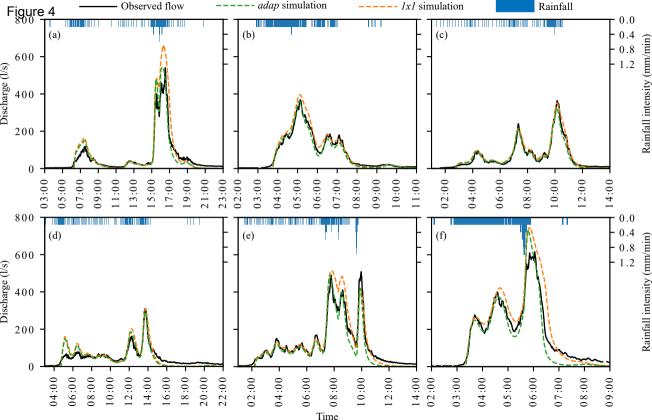


Fig. 1. Land cover and layout of the stormwater network in the Länsi-Pakila catchment. Surface runoff is routed to open storm sewer nodes, representing storm drain inlets and channel inlets, whereas runoff from connected roofs is routed both into open and closed storm sewer nodes, the latter representing manholes and pipe connections.

Fig. 2. SWMM computation times (min) for the calibration and validation events using *1x1* and *adap* models (Desktop PC, Intel Xeon 3.20 GHz CPU, Ubuntu Linux 16.04 LTS).

Fig. 3. Comparison of subcatchments and routing between (a) 1x1 and (b) *adap* models for the Länsi-Pakila catchment. The arrows depicting subcatchment routing are drawn between the subcatchment mass centers.

Fig. 4. Observed (5 min moving average) and *adap* and *1x1* simulated discharges for the calibration events (a) c1, (b) c2, and (c) c3 and for the validation events (d) v1, (e) v2, and (f) v3.